Properties of Indonesian Peat in Relation to the Chemistry of Carbon Emission

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Abstract: The discussion is focused on several properties of peat, such as the composition of peat materials, critical water content in relation to irreversible drying, concentration of -COOH, and the chemistry of C emission as a result of the reduction and oxidation processes of peat materials. The experiments were divided into two steps: (i) laboratory experiments and (ii) field experiments. For the laboratory experiments, peat samples of fibric, hemic and sapric decomposition degrees were taken from Jambi, Sumatra and Central Kalimantan. On the field, the experiments were focused in Central Kalimantan. Based on the research results, the composition of peat material is rather similar, being mostly rich in wood containing high content of lignin that varies from 65 to 80% for the peats of Jambi, and 78 to 93% for the peats of Central Kalimantan. The content of cellulose was mostly less than 10%, while hemi-cellulose and protein were not detected. The main organic acids, as a result of lignin biodegradation and the sources of C-release, are of aromatic group consisting mainly of derivative phenolic acids. The concentration of such organic acids ranging from the highest to the lowest is as follows: ferulic acid ≈ synapic acid > p-hydroxybenzoic acid > vanilic acid > syringic acid. The drying and wetting processes on the peat materials affected the stability of organic acids, which was indicated by the loss of C through CO₂ and CH₄ releases. The release of CO₂ and CH₄ from fibric peat was higher than that from hemic and sapric peats. To solve such problems, the addition of metal ion (Fe³⁺) as an ameliorant to the peat at a certain dosage was used. In the field experiments, the use of mineral soil and basic slag at the rate of 5% maximum Fe³⁺-sorption decreased the total of C-emission from peat of about 28 to 31%.

Key words: Indonesian peat, peat material composition, derivative phenolic acids, C-release.

1. Introduction

Recently, the peat-land utilization in Indonesia for agriculture has received much attention although the peat mostly contains: (i) poor to very poor in nutrients for plant growth and (ii) organic materials which are rather similar in composition, being very rich in wood, i.e. more or less decomposed trunks and branches derived from the former vegetation covers [Sabiham, 1988]. Several investigators have demonstrated that the composition of peats is related to the forest structure and the characteristics of vegetation cover. The use of peat forest for agricultural activities has led to widespread declines in organic carbon (C) and hence in peat quality. The declines occur because, in such activities, the loss of organic-C is not offset by the gains of C through the deposition of biomass.

The important physical properties of peat are high porosity [Sabiham, 2000] and very high content of water. Therefore, peat is naturally in anaerobic condition. Because peat materials have an irreversible drying property, however, the peat function as hydrologic controller would easily be disturbed. On the irreversible drying condition, particularly after a drainage system in peat land is constructed, a part of peat materials will not be able to reabsorb water. In fact, peat materials are the primary storehouse of organic-C. The loss of C through the reduction and oxidation of peat in the forms of CH₄ and CO₂ respectively is mostly emitted to the atmosphere. Yagi and Minami [1990] reported that the highest rate of CH₄ emission during cultivation period (44.8 g CH₄ m⁻²) was in rice field consisting of peat. The drainage of peat releases oxygen (O₂) into the surface, which promotes decomposition. It means that since peat deposits are a pile of carbon (C), they became potential materials of C-release that could be emitted to the atmosphere. The current annual average emission from Indonesian peat between 2000 and 2006 is estimated to be 903 MtCO₂ [BAPPENAS, 2009]. This condition shows that peat is very fragile and easily destroyed. The destruction rate will be faster if peat is poor in alkali and/or metal cations and below the peat layer there are old soil formations or sand (quartz) layers. In an anaerobic condition, the decomposition of organic matter in peat can actually occur even though it is at a very slow rate.

1.1. Nature of the Research Problem

The studies on the chemistry of C-emission in Indonesia have been given very little attention compared with the activities of peat utilization. Therefore, it could be understood that the reclamation of peat lands in the country has always been a matter of some controversies. Since the 1960s the Indonesian Government has started agricultural development on peat lands through the transmigration program, and from 1980s private companies started to open the lands for plantations. However, as a result of initial drainage, the stability of peat decreased and its land surface subsided due to the presence of high C-release from peat to atmosphere [WWF, 2008]. This means that the development of peat for agricultural activities has been hampered by the very little understanding of the people about the process of peat degradation.

Recognizing such problems particularly understanding of Indonesian peat properties in relation to the chemistry of C-emission has apparently become more important. The chemical composition of peat, critical water content, surface charge and adsorption characteristics of cation are the key factors that strongly affect the chemistry of C-emission. Low contents of water and cation in peat are also highly potential causes of C-losses.

1.2.Methodology Employed

The sites selected chosen for the study were the eastern part of Jambi, Sumatra and the center and southern parts of Central Kalimantan. The experiments, both in the laboratory and the experimental sites, were conducted during the period of 1999 until 2004. Peat deposits here are mostly situated inland of the coastal plains [Polak, 1975]. Therefore, peat samples were taken from the representative areas with different environments: marine, brackish and fresh water environments, which are characterized by a condition of water salinity in each environment. Peat land influenced by water inundation having water salinity of 1.0 to more than 2.0 mmhos cm⁻¹ is categorized into marine environment, and the land influenced by that having water salinity of 0.5 to 10 mmhos cm⁻¹ is categorized into brackish environment. The land classified into fresh water environment if it is influenced by water inundation having water salinity of less than 0.5 mmhos cm⁻¹.

The research was divided into two steps. The first (laboratory experiments) activities were focused on the inherent properties of peat, namely: (i) chemical compositions of peat in relation to the organic acids as the sources of C-release, for which organic acid measurement focused on derivative phenolic acids was conducted with the method of partition separation using reversed-phase column of C₁₈ (μBondapakTM 2.9x300 mm) and UV detector with D2-lamp at the wave of 280 nm; HPLC (High Performance Liquid Chromatography) was used to determine derivative phenolic acids, (ii) critical water content in relation to the process of irreversible drying determined by using the methods of Bisdom *et al.* [1993], (iii) adsorption characteristics of Fe³⁺ in peat, for which the methods of Fox and Kamprath [1971] modified by Widjaja-Adhi *et al.* [1990] and Syers *et al.* [1973] were used to select the cation that has high reactivity with organic acids, and (iv) measurement of CO₂ and CH₄ releases through reduction and oxidation processes. The second (field experiments) activities were focused on the interaction of organic acids and mineral materials containing high Fe³⁺ and its influence to rice production and C-emission.

2. Results and Discussion

2.1. Chemical Composition of Peat

Table 1 shows several physical and chemical characteristics of peats in Jambi and Central-Kalimantan. The degree of sapric, hemic and fibric decomposition was determined by using the data of fiber volume and pyrophosphate index [Lynn *et al.*, 1974]. There are several interesting results to discuss. Based on the ash content, the peats in Jambi and Central Kalimantan, which have the loss of ignition of mostly more than 95%, could be categorized as poor peats. C/N ratio of peats in both areas is still very high (more than 50%). The peat composition is dominated by lignin of 65% to 80% for the peats of Jambi, 78% to 93% for the peats of Central Kalimantan. The content of cellulose in all peat

samples was mostly less than 10%, while hemi-cellulose and protein were not detected. Lignin, as the dominant peat composition, is categorized as a thermoplastic system, highly aromatic polymers, derived from guaiacyl propane monomers [Tan, 1998]. Plant lignin in most Indonesian peats consists of three types of basic monomers, namely: (i) lignin from grasses and palm, in which the type of basic monomer is derived from 4-hydroxy phenyl propane, (ii) lignin from softwood derived from 3-methoxy-4-hydroxy phenyl propane, and (iii) lignin from hardwood derived from 3-5-dimethoxy-4-hydroxy phenyl propane as can be seen in Fig. 1. Orlov [1995] showed the processes of lignin disintegration that result in several derivative phenolic acids. In relation to plant growth, of 18 derivative phenolic acids as stated by Stevenson [1994], only five to six were the most important derivative phenolic acids found in the peats of Jambi and Central Kalimantan [Sabiham, 1997; Mario and Sabiham, 2002]: ferulic, synapic, *p*-coumaric, vanilic, syringic, and *p*-hydroxybenzoic acids (Fig.2).

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Peat properties	M	arine pea	at	Bra	ackish pe	eat	Fres	h water p	oeat
1 cat properties	S-ric	H-mic	F-ric	S-ric	H-mic	F-ric	S-ric	H-mic	F-ric
Jambi									
Organic-C (%)	54.9	56.9	-	55.8	56.5	57.4	54.8	56.0	55.8
Ash content (%)	5.26	2.08	-	5.62	3.53	3.73	3.75	2.68	1.09
Total-N (%)	1.06	1.04	-	0.82	0.94	0.78	0.86	0.76	0.71
C/N ratio	51.8	54.7	-	68.1	60.1	73.6	63.7	73.7	78.6
Lignin (%)	71.6	65.6	-	73.4	74.0	80.0	77.8	75.4	79.2
Cellulose (%)	5.9	9.6	-	6.84	5.4	3.4	5.2	2.6	4.8
Central Kalimantan									
Organic-C (%)	54.3	55.7	53.5	55.5	57.0	57.1	57.7	57.7	57.8
Ash content (%)	4.62	8.71	6.14	4.53	2.02	1.83	0.63	0.62	0.58
Total-N (%)	0.91	0.60	0.71	0.71	0.76	0.65	0.77	0.60	0.61
C/N ratio	59.7	92.8	75.4	78.2	75.0	87.8	74.9	96.2	94.8
Lignin (%)	78.8	86.7	82.8	87.5	91.5	90.5	93.0	92.0	92.2
Cellulose (%)	10.9	4.6	4.9	5.3	3.9	1.7	2.3	4.8	3.1

Notes: 1. In Jambi, marine fibric-peat could not be found due to intensive cultivation by local people.

- 2. S-ric = sapric peat; H-mic = hemic peat; F-ric = fibric peat.
- 3. Hemi-cellulose and protein contents in the peats of Jambi and Central Kalimantan were not detected.

Table 2 shows the concentrations of most important derivative phenolic acids found in the studied areas. Based on the concentration of the acids, ferulic and synapic acids are the highest, and the lowest was siringic acid. Such a series of the acid concentrations could be listed as follows: ferulic acid \approx synapic acid > *p*-coumaric acid > *p*-hydroxybenzoic acid > vanilic acid > syringic acid. These acids were categorized as the main sources of C-release due to the high content of carboxyl (-COOH) and methoxy (-OCH3) groups. COOH could be broken down completely into CO_2 and H_2O through the processes of oxidation-reduction. CO_2 could also be released when the methoxy groups ($-OCH_3$) changed into -OH groups during the phenol-OH formation through the processes of demethylation, hydroxylation and oxidation. The other important gas, methane (CH_4), is produced by strict anaerobic bacteria (methanogens) [van der Gon and Neue, 1995]. The rate at which anaerobic decomposition of peats proceeds as well as the resulted ratio of CO_2 to CH_4 particularly depended on soil pH. A major pathway in anaerobic condition produces CH_4 through the reduction of CO_2 with H_2 or organic molecules (H_2A) as the H donor if soil pH is very low [Neue and Scharpenseel, 1985], as follow:

$$CO_2 + 4H_2A \rightarrow CH_4 + 2H_2O + 4A$$

Although atmospheric CH₄ concentration is much lower than the atmospheric concentration of the most important greenhouse gas, CO₂, the rising atmospheric CH₄ concentration may significantly affect global temperature. The importance of CH₄ as a greenhouse gas is, for example, due to its relatively large annual concentration increase of approximately 0.8% [IPCC, 1992].

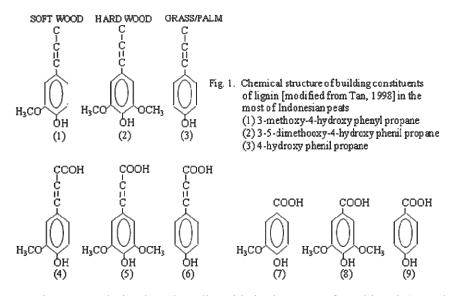


Fig. 2. The most important derivative phenolic acids in the peats of Jambi and Central Kalimantan: (4) Ferulic acid; (5) Synapic acid; (6) *p*-Coumaric acid; (7) Vanilic acid; (8) Syringic acid; and (9) *p*-Hydroxybenzoic acid.

Table 2. Concentrations of the most important derivative phenolic acids (mM) in the peats of Jambi and Central Kalimantan

Peat properties	M	arine pea	at	Bra	ackish pe	eat	Fres	h water j	oeat
1 cat properties	S-ric	H-mic	F-ric	S-ric	H-mic	F-ric	S-ric	H-mic	F-ric
Jambi									
<i>p</i> -Hydroxybenzoic acid	0.33	0.32	_	0.28	0.43	0.20	0.53	0.35	0.24
Vanilic acid	0.32	0.24	_	0.27	0.39	0.12	0.45	0.38	0.22
Syringic acid	0.31	0.24	_	0.40	0.29	0.15	0.37	0.30	0.13
<i>p</i> -Coumaric acid	0.33	0.53	-	0.53	0.35	0.18	0.70	0.45	0.32
Ferulic acid	1.10	0.83	-	1.21	0.29	0.10	1.10	0.80	0.52
Central Kalimantan	Com	posite sar	nple	Composite sample			Composite sample		
<i>p</i> -Hydroxybenzoic acid		0.65	•	0.80				0.91	•
Vanilic acid		0.66			0.78			0.84	
Syringic acid		0.58			0.71			0.84	
<i>p</i> -Coumaric acid		1.33		1.54			1.83		
Ferulic acid		1.42		1,65			2.24		
Synapic acid		1.35		1.76			2.32		

Notes: 1. In Jambi, marine fibric-peat could not be found, and synapic acid was not analyzed

- 2. S-ric = sapric peat; H-mic = hemic peat; F-ric = fibric peat.
- 3. In Central Kalimantan, peat samples were not separated between sapric, humic and fibric peats.

To determine the amount of COOH groups in peats, Ca acetate procedure [Tan, 1998] was used. The Ca acetate method made use of the reaction in which acetic acid (CH₃COOH) was formed and released according to the reaction as follows:

$$2R\text{-COOH} + Ca(CH_3COO)_2 \rightarrow (RCOO)_2Ca + 2CH_3COOH.$$

COOH content was determined by titration of the CH₃COOH solution with 1N NaOH, followed by calculation using the formula of:

COOH groups (me g^{-1} of peat) = $[(Vs - Vb) \times N]$ / weight of peat sample,

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where: $Vs = volume of CH_3COOH solution$; Vb = volume of blank; and N = normality of NaOH.

The analysis result of the COOH content can be seen in Table 3.

Table 3. The content of COOH functional groups (me g⁻¹ peat) in the peats of Jambi and Central Kalimantan

Decomposition		Peat of Jambi			Peat of Central Kalimantan			
degree	Marine	Brackish	F. water*	Marine	Brackish	F. water*		
Fibric peat	0.51	0.48	0.35	0.51	0.39	0.39		
Hemic peat	0.34	0.36	0.33	0.50	0.30	0.34		
Sapric peat	0.33	0.40	0.29	0.50	0.30	0.30		

Notes: *) F. water = Fresh water.

2.2. Critical Water Content and Irreversible Drying

The most important factor that influenced the COOH groups in peat was critical water content. The value of critical water content was calculated by making the relationship between the water content at certain levels and the proportion of irreversible drying.

Irreversible drying was determined by the method of "water drop penetration time". A model used for such relationship was $Y = ae^{-bX}$. To analyze the proportion of irreversible drying, peat samples were dried at 50 °C in an electric oven with time intervals of 15, 30, 45, 55, 65, 75, 90, 105, 135 and 150 minutes. The results showed that in all environmental conditions, namely marine, brackish and fresh water environments, fibric peat with the high values of critical water content (Table 4) needed a shorter period to reach an irreversible drying condition compared with hemic and sapric peats.

Table 4. Irreversible drying (ID) period and the value of critical water content (CWC) in the peats of Jambi and Central Kalimantan

Peat		Peat of Jambi		Peat or	f Central Kalir	nantan
properties	Marine	Brackish	F. water	Marine	Brackish	F. water
Fibric peat ID period (minute) CWC (%)**	90	90	90	90	90	90
	336.8–450.9	316.5–423.7	308.9–413.5	290.6–388.9	273.0–365.5	346.6–464.0
Hemic peat ID period (minute) CWC (%)**	120	120	105	120	105	105
	165.5–221.6	245.0–327.9	318.3–426.5	202.7–271.4	178.0–238.3	191.7–256.6
Sapric peat ID period (minute) CWC (%)**	135	105	120	135	135	120
	186.8–250.0	280.6–375.7	238.2–318.9	192.0–257.0	224.4–300.4	232.0–310.6

Notes: *) F.water = fresh water

In fact, the fibric peat that is categorized into a suborder of fibrist according to Soil Taxonomy [Soil Survey Staff, 1999] mainly consists of plant remains which are not totally destroyed yet. A part of botanical origin of peat materials could be readily recognized, and consists mostly of the partly decomposed wood with small amount of the plant remains of leave and grasses. As a consequence, colloidal organics in fibric peat are very small, so the peat could not absorb more water. On the other hand, the perfectly decomposed plant remains, such as sapric peats, produced more organic acids as colloidal organics which could tightly bind the water molecules, so that the water did not easily lose from its colloidal surface. Hemic peat consists of plant remains which have degrees of decomposition in between fibric and sapric peats.

^{**)} Below the critical water-content level, peat materials could not reabsorb water.

3.41

3.30

Based on Table 4 it can be mentioned that the higher the critical water content of peats, the more rapidly the peat materials would dry toward irreversible drying. Conversely, peat with low critical water content became very slow to dry. The high content of COOH groups (Table 3) might be one of the reasons why the above facts occurred. Tschapek *et al.* [1996] reported that functional groups of COOH are polarity and hydrophilic, so they can be helpful in the process of adsorbing water.

2.3. The adsorption characteristic of Fe³⁺

Sapric

Based on pH_{ZPC} and pH actual (Table 5), it can be mentioned that the surface of colloidal organic of peats mostly indicated as negative charge. In Indonesian peats which are mostly in acid condition, the charge is generally controlled by phenolic-OH and COOH groups. However, Tan [1998] noted that 54% of the exchange capacity of organic acids was attributed to COOH groups and they started to dissociate its H^+ at pH 3–5.5. This condition was very important for colloidal organic in the process of attracting cations, which lead to complex reaction (chelation). This fact was used to solve the problem in relation to the disintegration of organic matter through C-release.

1 1		•										
Decomposition		F	eats of	eats of Jambi				Peats of Central Kalimantan				
degree	Mar	rine	Brac	ckish	Fresh	water	Ma	rine	Brac	ckish	Fresh	water
	pH _Z *	pH_A	pH_Z	pH_A	pH_Z	pH_A	pH_Z	pH_A	pH_Z	pH_A	pH_Z	pH_A
Fibric	4.56	4.67	3.16	3.51	3.31	3.40	4.48	4.75	3.86	4.00	3.76	3.97
Hemic	4.13	2.48	3.24	3.47	3.28	3.35	4.28	4.33	3.43	3.78	3.51	3.02

3.33

3.54

4.53

4.82

3.15

3.52

Table 5. pH_{ZPC} and pH actual of the peats of Jambi and Central Kalimantan

5.12

3.36

Notes: *) pH_Z= pH_{ZPC} (pH at zero point of charge) is soil pH at which the surface charge of the colloidal organic is electrically neutral.

3.56

The use of selected cation of Fe³⁺ as an ameliorant for increasing peat stability was based on the previous research [Sabiham, 1997] in which the method of Langmuir curve according to Fox and Kamprath [1971, modified by Widjaja-Adhi *et al.*, 1990] and Syers *et al.* [1973] was adopted to select seven cations of Fe³⁺, Fe²⁺, Cu²⁺, Ca²⁺, Zn²⁺, Mn²⁺ and Al³⁺. Based on this research, the effectiveness of Fe³⁺ reaction with the colloidal organic to form the complexes is highest. Such a series of cation listed in the decreasing order of preferential interaction in the formation of complexes is as follows:

$$Fe^{3+} \approx Fe^{2+} > Al^{3+} > Cu^{2+} > Ca^{2+} > Zn^{2+} > Mn^{2+}$$
.

Maximum adsorption of Fe^{3+} in different pH increased, as presented in Table 6, with the increasing soil pH. Because the pH of peat is categorized as pH dependent charge, the increase in pH has caused the rise of negative charges due to the presence of functional group (COOH) dissociation, causing the increase in cation (Fe^{3+}) adsorption.

Table 6. The average of Fe³⁺-sorption (μg g⁻¹) in different soil pH of the peats in Jambi and Central Kalimantan

Location	Environ-			Soil pH		
	mental unit	-0.25 pH _{ZPC}	pH_{ZPC}	pH_{actual}	+0.5 pH _{ZPC}	+1.0 pH _{ZPC}
Jambi	Marine Brackish Fresh water	16,446.6 12,351.7 10,000.7	22,208.5 17,630.2 15,613.4	23,533.9 19,520.9 15,933.4	26,554.9 20.931.3 16,751.3	29,735.3 24.049.2 17,952.7
Central Kalimantan	Marine Brackish Fresh water	12,287.0 8,660.0 8,706.3	19,706.9 12,081.0 9,017.5	20,143.7 16,349.0 10,943.4	23,659.7 17,029.8 11,204.9	24,650.5 18,594.8 11,567.9

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2.4. Measurement of C-release

It has been known for a century that agricultural activities on peats lead to the subsidence of peat surface and C-loss as CO₂ and CH₄ emissions. The loss has mainly been considered as an agricultural problem, especially when the peat is overlying mineral soil that is unsuitable for cultivation. Recently, the effect of cultivation of peats for agriculture in Indonesia, particularly for plantation, has caused the increase in C-releases, and it has become a major concern. However, direct CO₂ measurements in Indonesian peats are generally limited to few locations and only conducted during a very short period of time. Few solutions to such problems were also found.

CO₂ efflux from Indonesian cultivated peats estimated by BBSDLP (Agricultural Land Resource Institute) [2009] reached up to 73 ton CO₂e ha⁻¹ year⁻¹ from plantation areas, 27 ton CO₂e ha⁻¹ year⁻¹ from annual crop cultivation, and 55 ton CO₂e ha⁻¹ year⁻¹ from annual fire. An assessment of Indonesian peat GHG emission from peat oxidation according to IPCC Tier 2 standards using land use and land cover data from 2000-2006 showed the average annual net emissions of 220 Mt CO₂ yr⁻¹ [BAPPENAS, 2009]. Although C-emissions from peats are potentially large, the uncertainty of these estimates is also quite large.

The present paper shows measurements of C-release from the cultivated peat starting from laboratory experiments and followed by field experiments, and the results give a suggestion how to solve the problems of C-release practically. To measure the CO₂ and CH₄ releases resulted from the decomposition process, the method of Chapman *et al.* [1996] was used. An incubation process of peat samples was done in a glass bottle of 165 ml within three weeks. The samples were put in the glass up to 2.2 cm thick. The instruments used for this experiment were Gas Chromatography Shimadzu 14-B and Chromatopac Shimadzu C-R6A. The research results showed that the decomposition process of peats had a great effect on C-release. Table 7 shows C-release in the forms of CO₂ and CH₄ fluxes. Fibric peat produced the highest of C-fluxes compared with hemic and sapric peats. The use of Fe³⁺ as ameliorant was able to reduce CO₂ and CH₄ from all peat degrees of decomposition (Table 8).

Table 7. CO₂ and CH₄ fluxes (μg g⁻¹ h⁻¹) from the peats of Jambi and Central Kalimantan

Forms of		Peats o	f Jambi		Peats of			
C- fux	Fibric	Hemic	Sapric	Average	Fibric	Hemic	Sapric	Average
CO_2	53.8	49.4	45.0	49.4	65.1	58.8	53.5	59.1
$\mathrm{CH_4}$	29.5	14.9	11.8	22.7	22.5	19.4	16.1	19.3

Based on the contents of CO_2 , CH_4 and Fe-solution after giving Fe^{3+} as the ameliorant (Table 8), the rate of 5% maximum sorption of Fe^{3+} was adopted as the basic treatments for field experiments. Although the use of Fe^{3+} at the rate of 7.5% maximum sorption sharply decreased the CO_2 and CH_4 productions of about 42.1 to 47.7% and 30.6 to 34.8%, respectively, but Fe-solution is higher than that resulted by the use of Fe^{3+} at the rate of 5% maximum sorption and the CO_2 and CH_4 production is not much different. The decrease is caused by the formation of stable complex binding between Fe^{3+} and organic acids as can be shown by the illustration of the reaction as follows (Fig.3):

Vanilic acid

$$Fe^{2+} + 3/2 O_2$$
 $Fe^{3+} + CO_2 + 2H_2O_3$

Complex reaction / chelation (Stable complex binding)

Fig. 3. An example of interaction between vanilic acid and Fe²⁺ in a condition of low soil-pH

With such formation, peat would be more resistant to decomposition process, thereby reducing the production of CO₂ and CH₄.

Table 8. Effect of the application of Fe³⁺ on the CO₂ and CH₄ productions (μg g⁻¹ h⁻¹) and the soluble Fe (ppm) in composite-peat samples from Jambi and Central Kalimantan

Rate of Fe ³⁺ *	P	eats of Jamb	i	Peats of	Central Kal	imantan
(% max. sorption)	CO_2	CH ₄	Fe-solution	CO_2	CH ₄	Fe-solution
2.5	36.2	17.8	16.3	37.0	16.7	16.2
5.0	30.9	15.2	18.5	30.5	14.1	18.9
7.5	28.6	14.8	26.3	28.0	13.4	27.6

Notes: *) Given in the form of FeCl₃.6H₂O

2.5. Interaction of organic acids and mineral materials containing high Fe³⁺, and its influence on rice production and C-emission

It has long been my contention that the production of agricultural crops planted on peats is always lower than those planted on mineral soils. The low productivity of peats is mainly caused by the constraint of high content of organic acids, particularly derivative phenolic acids which are phytotoxic to plants. Peats are also very easy to get destroyed due to the high loss of organic-C upon reclamation. Rieley *et al.* [1997] stated that peats in Southeast Asia are formerly forests and are of only marginal use for growing crops, yet with the conversion to agriculture together with non-sustainable logging, are a major and increasing threat.

Recently, however, many scientists have conducted research in cultivated peatlands in order to solve the constraints. In the present paper, I argue that a method to decrease the derivative phenolic acids by changing the environment from anaerobic into aerobic conditions should be considered. The addition of mineral soil enriched with basic slag containing high levels of Fe^{3+} is expected to be able to reduce the harmful effects of such acids, meaning that the productivity and stability of peats can be increased. The experiments were carried out in three locations of Central Kalimantan's peatlands: marine peatland located in Samuda, brackish peatland in Sampit, and fresh water peatland in Berengbengkel. A combination of mineral soil containing Fe^{3+} of about 22.06% of Fe_2O_3 and basic slag of about 42.6% of Fe_2O_3 on various levels was given to the peat as ameliorants based on 5% of the maximum adsorption of Fe^{3+} . The combinations of such ameliorants were MBS_0 = without treatment; MBS_1 = 100% mineral soil (ms), MBS_2 = 50% ms + 50% slag (bs); MBS_3 = 100% bs.

The research results showed that the concentration of observed derivative-phenolic acids, namely: ferulic, synapic, *p*-coumaric, vanilic, syringic and *p*-hyroxybenzoic acids in peats decreased with the addition of mineral soil or basic slag, or the combination of both materials (Table 9). The higher the mineral soil in peats, the more interaction would actually emerge between organic and mineral materials containing high Fe³⁺. Sabiham [1997] stated that the proportion of Fe³⁺ in peats is mainly distributed in the forms of chelation and strong bond which varies from 55.1 to 78.6% of the total extracted Fe³⁺.

The addition of ameliorants also increased the rice production, as can be seem in Table 10. Based on the observation results, however, rice did not seems to be able to grow well on fresh water peat; without ameliorant, plant growth was totally hampered, so after 42 to 49 days of transplanting the rice died. Tadano *et al.* [1992] reported that derivative phenolic acids, such as ferulic, synapic, p-coumaric, and p-hydroxybenzoic acids are phytotoxic for rice, particularly during the first stage of plant growth. He also mentioned that ferulic acid in peat is more toxic compared to the other derivative phenolic acids. Tsutsuki *et al.* [1994] stated that the concentration of phenolic acids at the range of 0.6 to 3.0 mM could hamper the root growth of rice up to 50%.

Based on Table 10, the production of rice in marine peat is higher compared to that of rice in brackish and fresh water peats. Marine peat is categorized as eutrophic peat [Andriesse, 1988], rich in nutrients due to very close to the marine environment, so the peat could support plant growth well. On the other hand fresh water peat categorized as oligotrophic peat, in which the source of nutrients is

mostly derived from rain water, does not seem to be easy to bring into agricultural activities. Therefore, it could be suggested that the natural fresh-water peat not converted into agricultural use, as it is extremely acid and infertile.

Table 9. Average of the observed derivative-phenolic acids in peat of Central Kalimantan treated by some combinations of mineral soil and basic slag containing high content of Fe³⁺

No	Treatment		De	erivative phen	olic acid (m	M)	
INO	Heatment	Ferulic	Synapic	<i>p</i> -coumaric	Vanilic	Syringic	p-HB*
1. N	Iarine peat						
	MBS_0^***	1.43	1.35	1.33	0.66	0.58	0.65
	MBS_1	0.76	0.91	0.91	0.43	0.35	0.49
	MBS_2	0.92	0.97	1.06	0.51	0.43	0.51
	MBS_3	0.99	1.02	1.10	0.53	0.44	0.52
2. B	Brackish peat						
	MBS_0	1.65	1.76	1.54	0.78	0.71	0.80
	MBS_1	0.93	1.19	0.95	0.55	0.48	0.58
	MBS_2	1.21	1.30	1.20	0.63	0.59	0.65
	MBS_3	1.22	1.40	1.17	0.65	0.60	0.67
3. F	resh water peat						
	MBS_0	2.25	2.32	1.83	0.84	0.84	0.90
	MBS_1	1.51	1.49	1.46	0.64	0.61	0.66
	MBS_2	1.70	1.73	1.56	0.68	0.70	0.72
	MBS_3	1.81	1.82	1.59	0.69	0.73	0.72

Notes: *) p-HB = p-hydroxybenzoic

Table 10. The effect of ameliorants on the weight of rice grain of IR-64 (t ha⁻¹)

Treatment	Peats of Central Kalimantan						
Treatment	Marine peat	Brackish peat	Fresh water peat				
${ m MBS_0}^*$	4.09	2.12	0.00**				
MBS_1	4.18	2.27	0.00				
MBS_2	5.17	4.78	1.15				
MBS_3	5.62	5.08	1.56				

Notes: *) MBS = combination of mineral soil and basic slag as ameliorant: MBS₀ = without ameliorant; MBS₁ = 100% mineral soil (ms); MBS₂ = 50% + 50% basic clag (bs); MBS₃ = 100% bs.

Table 10 shows that although the treatment of MBS_3 (100% basic slag or bs) produced the highest level of yield, the production resulting from the treatment of MBS_2 at the same environment is not so much different. Based on the practical aspect, it can be suggested that the treatment of MBS_2 (50% ms + 50% bs) is more useful for improving the peat productivity.

In fact, undisturbed naturally forested peatlands either have a balanced C-budget or show a net accumulation of C through the natural process of peat formation. Therefore, peat in natural condition can be categorized as a stable peat, meaning that the accumulation of C is higher than its loss through decomposition process. The C-sequestration rate from natural peatlands in Indonesia has been estimated to be up to 0.8 t C ha⁻¹ yr⁻¹ [Page *et al.*, 2004]. In cultivated peatlands, however, most agricultural crops grown there have always required drainage of peat, even though the drainage releases oxygen into the surface which promotes decomposition. As a consequence, the peat stability is disturbed, and C-release from the peat increases.

^{**)} MBS = combination of mineral soil and basic slag as ameliorant: MBS₀ = without ameliorant; MBS₁ = 100% mineral soil (ms); MBS₂ = 50% + 50% basic clag (bs); MBS₃ = 100% bs.

^{**)} Rice plant died at 42 - 49 days after transplanting.

To measure the flux of CO₂ and CH₄ in the field, a chamber for trapping the gases made from the fiberglass with the size of 1 m x 0.5 m x 0.5 m, was used. Syringes were used to take the samples of gases from the chamber. The samples were then put on the vacuum bottles. In this research, Gas Chromatography Shimadzu 14-B and Chromatopac Shimadzu C-R6A were used to determine the CO₂ and CH₄ emissions. The emissions were calculated by using the following equation [Boer *et al.* 1996]:

Based on the research results, the average of total C-losses through CH_4 and CO_2 emission from the peats in Central Kalimantan where IR-64 is continuously planted in submerged condition is estimated be about 2.09 t C ha⁻¹ yr⁻¹ in fresh water peat, 1.99 t C ha⁻¹ yr⁻¹ in brackish peat, and 1.97 t C ha⁻¹ yr⁻¹ in marine peat. However, by using the ameliorants of (50% ms and 50% bs) the total C-loss decreased to about 1.49 t C ha⁻¹ yr⁻¹ (28%) in fresh water peat, 1.38 t C ha⁻¹ yr⁻¹ (30%) in brackish peat, and 1.34 t C ha⁻¹ yr⁻¹ (31%) in marine peat. In fact, these calculation results of C-emissions are still lower than those reported by Maltby [1997], i.e. up to 5-42 t C ha⁻¹ yr⁻¹.

3. Conclusion

The Indonesian peats, represented by the peats of Jambi and Central Kalimantan, are mostly easy to get destroyed, meaning that the stability of such peats is very low. The low stability is characterized by the C-release in the form of CH₄ and CO₂ emissions as a result of the decomposition process. To maintain the stability of peats, water content should always be kept higher than the critical water content. Peats treated by ameliorants containing high Fe³⁺ in submerged condition are relatively resistant to decomposition. By using the ameliorants at the rates (50% ms + 50% bs) as the basis of 5% maximum sorption Fe³⁺, derivative phenolic acids as the dominant organic acids in peats and as the sources of C-emission would decrease and peat productivity would increase. Cation metal of Fe³⁺ has potential as Cation Bridge to enhance the polymerization of derivative phenolic acids by linking the individual phenolic molecules together to yield a chain-like structure, and it results in the increase of peat material stability. The decrease of C-emission is 0.60 t C ha⁻¹ yr⁻¹ (28%) in fresh water peat, 0.61 t C ha⁻¹ yr⁻¹ (30%) in brackish peat, and 0.63 t C ha⁻¹ yr⁻¹ (31%) in marine peat.

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